

APPENDIX AA

Paleontological Resource Assessment

October 13, 2008

David Sawyer
Analytical Environmental Services
1801 7th Street, Suite 100
Sacramento, CA 95811

Paleontological Sensitivity Assessment Report for the Madera and North Fork Sites

I. Project Description

A paleontological resource assessment for the North Fork Rancheria project was conducted on two potential sites under investigation. The North Fork Site is located approximately 40 miles to the east in the Sierra Nevada foothills, near the town of North Fork, California. The Madera Site is located west of and adjacent to Highway 99, north of the City of Madera, California. The North Fork Rancheria is proposing to build a casino and/or hotel development on one of these two locations.

The objectives of this assessment were to search and review literature and records on information pertinent to both sites, and to identify significant geologic formations and associated vertebrate fossils. Field surveys were conducted on both sites to examine the surface environment and assess the potential presence of paleontological resources. Such information will affect recommendations for mitigation plans deemed necessary for monitoring, salvaging, and preservation of vertebrate fossils for future scientific research. This assessment was conducted to assist Analytical Environmental Services with compliance responsibilities under CEQA and NEPA.

II. Findings (Appendices A-E and Plates I-V)

i) North Fork Site

Bedrock at the North Fork Site is entirely Mesozoic granitic igneous rock that is devoid of fossils. Paleontological repository locality records indicate that there are no vertebrate findings at or in the vicinity of the proposed site, and no fossils were encountered during the field survey.

ii) Madera Site

Geologic maps and literature on the Madera site area indicates Quaternary alluvial sediment is present throughout the study area. Field reconnaissance indicates surficial deposits consist of graded and tilled topsoil with areas of hardpan. Mid- to-Late Pleistocene deposits below this topsoil consist of three stratigraphic units from top to bottom: Modesto Formation, Riverbank Formation, and Turlock Lake Formation. Adjacent sites have produced vertebrate fossils in all three units (Dundas et al., 1996; Hilton et al., 2000). The Fairmead Landfill locality (UCMP V93128), located west of Hwy 99 and approximately 6 miles north of the study area, has yielded the largest deposit of Pleistocene vertebrates in the San Joaquin Valley. Vertebrate fossils were discovered in May 1993 during the excavation of a 5-acre

expansion cell. Initial studies of the site were conducted by the University of California Museum of Paleontology with further recommendations to monitor and salvage fossils exposed during excavations. Preliminary observations and results are presented in Dundas et al. (1996) and Dundas and Blades (1999). An updated list of taxa is provided in Kottachchi et al., 2008. Monitoring continues today under Lead Paleontological Monitor, Niranjala Kottachchi, and is expected to continue for 20+ years.

To date, over 5000 fossil specimens have been recovered from an area of 15 acres and depths of four to 20 meters below the surface. The majority of the vertebrates are from the upper unit of the Turlock Lake Formation from depths of six to eight meters below the surface in the west to depths of four to five meters below the surface to the east. Although no visible surface exposures of the Turlock Lake Formation exist in the Madera Site study area, it is possible that the fossil-bearing unit will be encountered at depth.

The main unit at the Madera Site is the Riverbank Formation (Marchand (1976), Marchand and Allwardt (1981)). Vertebrate fossils have been recovered from the middle unit of the Riverbank Formation at Fairmead Landfill as well as at other localities (Dundas et al., 1996; Hilton et al., 2000) and therefore, it is likely that fossils are present at the Madera Site. Although the Modesto Formation is absent at the Fairmead Landfill, geologic maps indicate it is present at the Madera Site. Repository locality records indicate only three other sites further north where single or few Pleistocene vertebrate fossils have been recovered from this stratigraphic unit. Therefore, this unit, where present in the study area, should be approached with caution.

III. Recommendations

Salvaging and preservation of paleontological resources have significant scientific and educational value. Monitoring paleontologically rich sites during excavation reduces the adverse impact on these valuable resources. Since the North Fork site is underlain entirely by igneous rock devoid of fossils, monitoring of the site is not necessary. However, Pleistocene vertebrate fossils are probably present in units underlying the Madera Site so all excavations associated with unearthing of *in situ* sediment below one to two meters should be monitored. Should fossil resources be encountered, the contractor must submit a Paleontological Resource Impact Mitigation Plan (PRIMP) outlining in detail the procedures for collecting (i.e. geographic and stratigraphic information) and preserving the fossils (i.e. stabilization methods). All fossils recovered during mitigation should be accessioned in an accredited scientific institution, such as the University of California Museum of Paleontology. Upon completing all monitoring, salvaging, and fossil preparation, the contractor must submit a final report detailing the results of the mitigation program.

Sincerely,
Niranjala Kottachchi, M.Sc.
Paleontologist



BRADY and ASSOCIATES

Geological Services

California Professional Geologist, Oregon Registered Geologist

Stream Restoration • Engineering Geology • Geophysics • Exploration Geology
2703 N. Adoline Ave., Fresno CA 93705 • (559) 284-1501 • Brady_geology@msn.com

Oct. 8, 2008

Mr. David Sawyer
Analytical Environmental Services
1801 7th Street, Suite 100
Sacramento, CA 95811

RE: Technical review

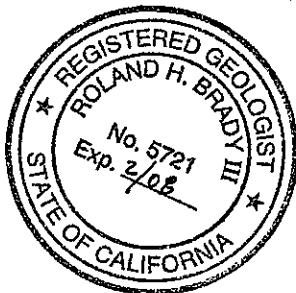
I have reviewed the content of your report entitled "Paleontological Sensitivity Assessment Report for the Madera and North Fork Sites" and its appendices. This work was written for the North Fork Rancheria and describes the potential for vertebrate resources to be present at proposed construction sites in California.

I found the methods, resources used, descriptions, and conclusions used therein to be consistent with current, professional geological standards.

Sincerely,

Roland H. Brady III

Roland H. Brady III, Ph.D.
California Professional Geologist #5721



received
10/17/08

The Fairmead Landfill Locality (Pleistocene, Irvingtonian), Madera County, California: preliminary report and significance

Robert G. Dundas¹, Randall B. Smith² and Kenneth L. Verosub³

¹Research Associate, Museum of Paleontology, University of California, Berkeley, CA 94720; Present address: Department of Anthropology, University of Montana, Missoula, MT 59812

²Department of Geology, California State University, Fresno, CA 93740-0024

³Department of Geology, University of California, Davis, CA 95616

ABSTRACT

A diverse vertebrate fauna, dominated by large herbivorous mammals, was discovered in May of 1993 at the Madera County Fairmead Landfill in alluvial fan, fan channel and marsh/lacustrine sediments representing the upper unit of the Turlock Lake Formation. Taxa identified from this fauna include: *Clemmys marmorata*, Anatidae, *Glossotherium harlani*, *Nothrotheriops* cf. *N. shastensis*, *Megalonyx* sp., *Canis armbrusteri*, *Canis* cf. *C. latrans*, *Smilodon* cf. *S. fatalis*, *Homotherium* sp., *Thomomys* sp., cf. *Dipodomys* sp., *Lepus* sp., *Mammuthus columbi*, *Equus* sp. (large and small), *Camelops* sp., *Hemiauchenia* sp., *Tetrameryx irvingtonensis*, *Capromeryx* sp., and *Odocoileus* sp. A late Irvingtonian age is indicated for the fauna based largely on the presence of *Tetrameryx irvingtonensis* coupled with the absence of *Bison*. The fossil bearing stratum is normally magnetized, and is inferred to have been deposited during the Brunhes normal magnetic polarity chron, placing an upper bound on the age of the fauna at 780,000 years before present (yr. B.P.). The Fairmead Landfill yields the first diverse latest Irvingtonian fauna from north-central California. Comparison to the older, coastal type Irvington fauna will enhance our understanding of the Irvingtonian of California, permitting better comparisons and correlations with other North American Irvingtonian age faunas.

INTRODUCTION

The Pleistocene of North America is divided into two land mammal ages (NALMAs), the Irvingtonian and the Rancholabrean, typified by California localities. Irvingtonian (early to middle Pleistocene) localities are sparse, in comparison to Rancholabrean (late Pleistocene) sites, and the precise stratigraphic position of many localities within the Irvingtonian is uncertain (Kurtén and Anderson, 1980). The type Irvington fauna occurs in reversely magnetized strata referred to the Matuyama magnetic chron, 780,000+ yr. B.P. (years before present; Lindsay et al., 1975; Kurtén and Anderson, 1980; Baksi et al., 1992). The exact position of Irvington within the Matuyama is questionable, but faunal constituents indicate that it most likely is in the upper portion of the magnetic chron. Until 1993, Irvington was the only diverse Irvingtonian age fauna known from north-central California.

A new fauna from an expansion cell of a landfill south of Chowchilla, Madera County, California augments this limited record. The Fairmead Landfill fauna, occurring in normally magnetized strata referable to the lower Brunhes magnetic epoch, is younger than the type Irvington fauna. Together the Irvington and Fairmead Landfill sites provide a record of the middle to late Irvingtonian of central

California, allowing for better comparisons and correlations with other North American Irvingtonian faunas.

To advance our understanding of the Irvingtonian in California a comprehensive study of the newly discovered Fairmead Landfill fauna and a revision of the type Irvington fauna are underway. The latter effort is necessary in order to report significant unpublished data (e.g. Firby, 1968), including many taxonomic records not cited in the literature. As an initial step in this process, some preliminary findings of the Fairmead Landfill locality study are presented here.

SITE LOCATION

The Madera County Fairmead Landfill, located at 21739 Road 19, Chowchilla, California, is about 3.3 km south-southeast of the junction of California state highways 99 and 152. Map coordinates are 37°03'24" N, 120°11'45" W, Berenda, California, United States Geological Survey 7.5" topographic quadrangle, 1987 photorevised (Figure 1). The fossil material was recovered at elevations of about 61-64 meters, which is 11-14 meters below ground surface. Fairmead Landfill is recorded as University of California Museum of Paleontology (UCMP) vertebrate fossil locality V93128.

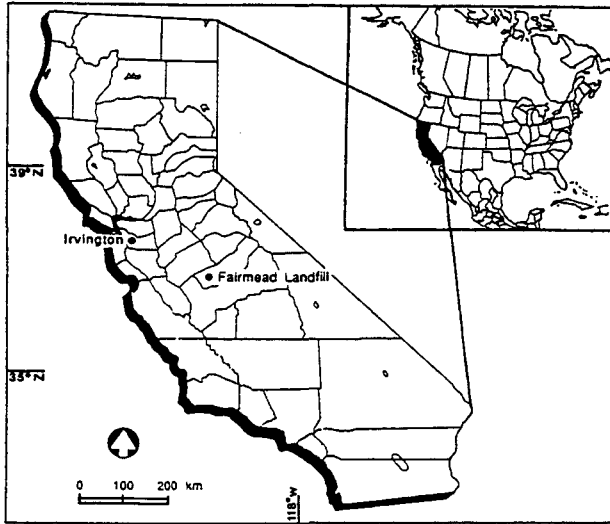


Figure 1. Map noting the locations of Fairmead Landfill, Madera County and Irvington, Alameda County, California. Base map is modified from Fay and Thiessen (1993).

HISTORY OF INVESTIGATIONS

Vertebrate fossils were discovered in May of 1993 during excavation of an expansion cell at the Madera County Fairmead Landfill. Following the initial evaluation of the site, the University of California Museum of Paleontology entered into a short-term, paleontological mitigation contract with Madera County to salvage the exposed vertebrate remains. A five meter square grid system, with alphanumeric coordinates, was used to map the position of excavated fossils. Recovery work began June 7, 1993 and concluded December 31, 1993, yielding thousands of specimens. Fossil recovery by the UCMP ended at Fairmead Landfill in the fall of 1993 following expiration of the mitigation contract between the UCMP and Madera County. Since the fall of 1993 R. G. Dundas has functioned as the professional vertebrate paleontologist monitoring the site to assure compliance with CEQA regulations regarding any future discoveries of fossil material at the landfill. D. L. Blades was hired by Madera Disposal Systems and the County of Madera in the fall of 1993 to act as on-site monitor under the direction of R. G. Dundas. Further landfill expansion from early 1994 to the present time has resulted in intermittent finds of bone. Fossils are removed as they are uncovered during ongoing landfill excavation. The UCMP is the repository for all fossil material from Fairmead Landfill.

GEOLOGY

The Fairmead Landfill is situated on the alluvial fan of the Chowchilla River, about 19 kilome-

ters west of the fan head. The 1993 expansion pit was excavated to a depth of about 14 meters through a sequence of sand, clayey sand, and minor clay which was well-exposed in temporary vertical cuts prior to final beveling of the pit walls. This sequence of distal alluvial fan and fan channel deposits is divisible into three stratigraphic units separated by unconformities and prominent paleosols (Figure 2). These units are designated A, B, and C from the surface downward. The vertebrate remains are from the lower exposed extent of Unit C.

Unit A consists of three to four meters of sand and pebbly sand with small-scale cross-bedding, representing the sandy fill of a former fan channel. The soil developed on Unit A is a yellow-brown sandy loam with a weakly to strongly-cemented iron-silica duripan, characteristic of the Madera soil series of the San Joaquin Valley.

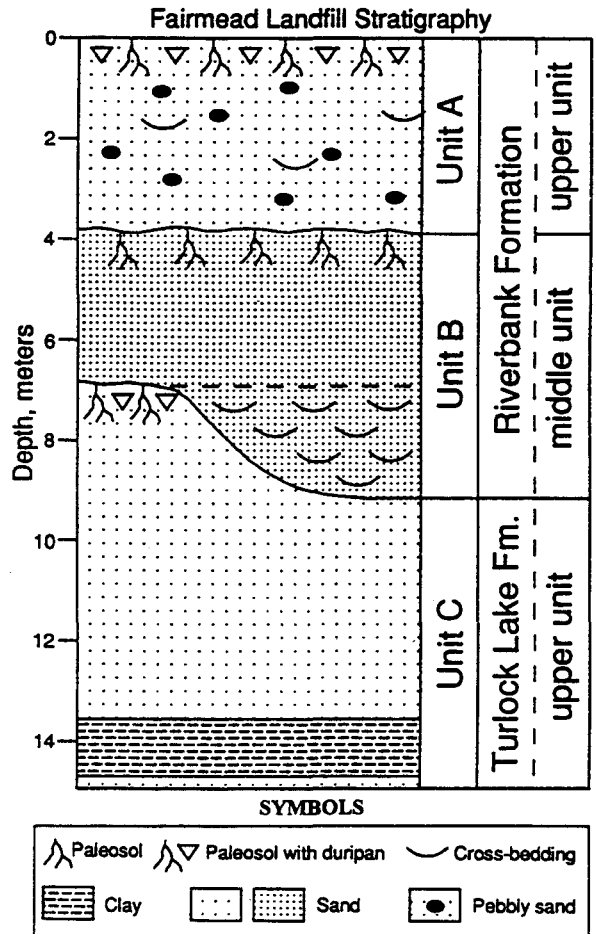


Figure 2. Generalized stratigraphy of the Fairmead Landfill 1993 expansion pit, based on south wall exposures and trenches in the pit floor. Also shown are inferred correlations between locally-defined units (A, B, and C) and the regional Pleistocene stratigraphy of the northeast San Joaquin Valley (right-hand column).

Unit B consists of three to four meters of massive clayey sand which in the western half of the pit overlies a sand-filled channel eroded several meters deeper into the underlying Unit C. The channel-fill consists of medium to coarse sand with pervasive trough cross-bedding, which indicates paleoflow toward the southwest and west. Although the western margin of this channel was not revealed by the 1993 excavation, its dimensions are probably comparable to those of the modern Chowchilla River. The top of Unit B is marked by a clay-rich paleosol overlain by a sharp erosional contact with the clean, unconsolidated sand at the base of Unit A. The Unit B paleosol has a blocky to prismatic structure with pervasive oxide staining of ped surfaces, but lacks a duripan. Pedogenic features persist to a depth of 1.3 m below the top of the unit.

Unit C comprises the lower half of the pit wall on the east side, with a maximum exposed thickness of six to seven meters. Most of the section consists of clayey sands and sands overlying a layer of massive sandy clay and clayey very fine sand which forms most of the pit floor. Abundant vertebrate fossils were found in the basal sandy clay and in the lowermost sand beds (Figure 3).

The sand layers in the upper part of Unit C are tabular and in most cases massive and clayey; layers of clean fine sand with cross-lamination occur locally. The coarser layers (medium to coarse sand) have sharp bases and show normal grading in their upper parts. Rootlet traces are very common in the upper few centimeters of most beds. These features point to rapid but intermittent deposition beyond the confines of active fan channels, probably during overbank flood and sheetflood events.

The sandy clay layer at the pit floor yielded most of the large vertebrate remains. Backhoe trenches revealed that the clay is as much as a meter thick, is cut by pervasive fine rootlet traces, and locally overlies a coarse pebbly sand. *Clemmys marmorata* and Anatidae remains found in the clay indicate the presence of a perennially wet environment, such as a marshy, abandoned fan channel or lake-margin. Many smaller bones in the clay show evidence of breakage *in situ*, and a few are oriented at high angles to the horizontal, both probably as a result of trampling by large herbivores. Floating pebbles in the clay surrounding the trampled bones were most likely mixed upward into the mud from the underlying pebbly sand.

A well-developed paleosol is present at the top of Unit C in the eastern part of the pit, but is absent in the western half owing to erosional downcutting of the channel at the base of Unit B.

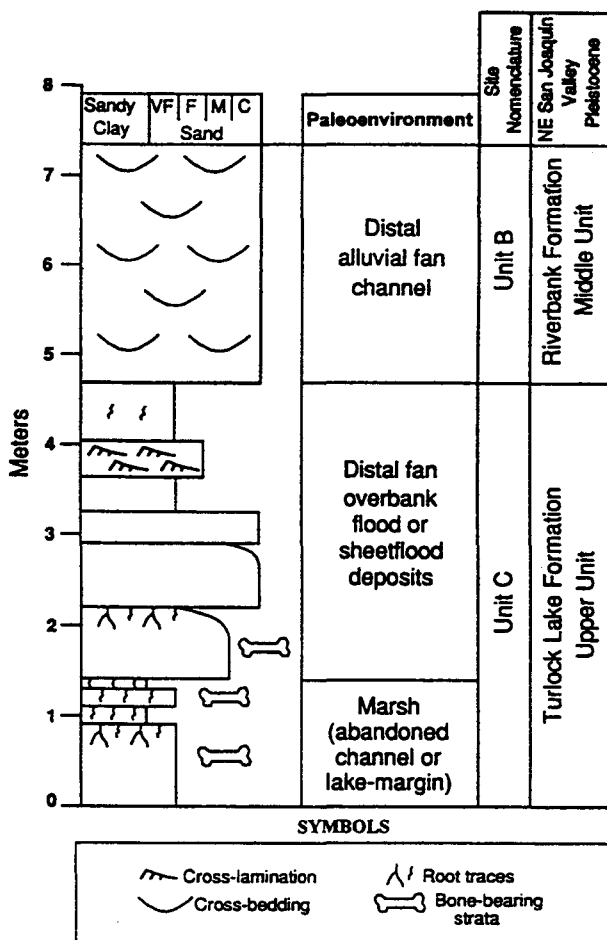


Figure 3. Detailed lithologic section for the lower part of the south wall of the Fairmead Landfill 1993 expansion pit between grid lines 11 and 12, showing inferred paleoenvironments and regional correlations.

The paleosol is clayey and very red in fresh exposures, with a strongly-cemented duripan.

Marchand and Allwardt (1981) mapped and subdivided the Pleistocene alluvial deposits of the northeastern San Joaquin Valley (Figure 4). The deposits exposed on the distal parts of the alluvial fans (such as in the Fairmead area) are subdivided on the basis of differing soil profile development and unconformities associated with buried soils. Their mapping in the area of the Fairmead landfill assigns the surface deposits to the middle unit of the Riverbank Formation, with the upper unit of the Riverbank Formation onlapping from the west and filling channels incised into the middle unit. Over most of the Chowchilla fan, the lower unit of the Riverbank Formation is missing, and the middle unit lies directly on the upper unit of the Turlock Lake Formation, which is locally exposed along incised channels beginning about 5 kilometers northeast of the landfill. In mapping the Pleis-

tocene deposits along the Chowchilla River, Helley (1966) also found that the Riverbank Formation is very thin on the fan east of Chowchilla, but thickens rapidly westward.

Preliminary correlations between the Fairmead sequence and the regional stratigraphy of Marchand and Allwardt (1981) are presented in Figures 2 and 3. We constructed contours on the exposed Riverbank-Turlock Lake contact northeast of Fairmead and used the resulting gradient to project that contact southwestward into the landfill area. Although this procedure is crude, it suggests that the fossil-bearing Unit C, exposed in the lower walls of the Fairmead pit, correlates with the upper unit of the Turlock Lake Formation. The overall thickness of Units A and B (7 to 9 m) at Fairmead accords well with the total thickness of the Riverbank Formation (3 to 7 m) measured by Helley (1966) 6.5 kilometers northwest at Chowchilla, which occupies a similar position on the Chowchilla fan. We suggest that Unit A is a previously unmapped Upper Riverbank channel fill, and corre-

late Unit B with the middle unit of the Riverbank Formation.

PALEOMAGNETISM

Samples were collected for paleomagnetic analysis from eighteen horizons in Unit C. Each sample was collected in a small (1.8 cm x 2.0 cm x 2.0 cm) plastic box that was placed over a pedestal of sediment carved on the outcrop. Each box was fully oriented with respect to a geographic coordinate system. The samples were collected at two sites in the Fairmead landfill. The first site was located on the southwest corner of the excavation in the clay layer at the basal exposure of the fossil-bearing bed. There were eight sampling horizons at this site, comprising three groups. The first horizon of the first group represented the lowest exposed level of the clay layer. A second horizon was located 32 cm above the first. The second group of three more horizons, separated by intervals of 16 cm and 14 cm respectively, was located about 0.5 m above the first group. A third group of horizons was located another 0.5 m above the second group. Within this group, the horizons were separated by 18 cm and 13 cm, respectively.

The second site was located on the south wall of the pit. There were ten sampling horizons at this site, and again the first horizon represented the lowest level exposed on the wall. A second horizon was located 30 cm above the first and a third horizon was 25 cm above the second. Another group of three horizons was located about 1.1 m above the highest horizon of the first group. These three horizons were separated by distances of 30 cm and 25 cm, respectively. A third group of two horizons was located another 1.1 m above the second. The two horizons in this group were separated by 30 cm. The final group of horizons began about 2.2 m above the third group. The two horizons in this group were separated by 27 cm. Together the two sites provide a comprehensive sampling of Unit C from its lowest exposed level in the landfill to within 1 meter of its contact with Unit B.

One sample from each horizon was analyzed in the Paleomagnetism Laboratory of the University of California at Davis. Paleomagnetic measurements were made using a 2-G Enterprises Model 760 cryogenic magnetometer. Alternating field demagnetization was done using a Schonstedt SSM-1 demagnetizer. Each sample was demagnetized at 10 different levels of peak alternating magnetic field: 0, 5, 10, 15, 20, 25, 30, 40, 50 and 60 mT. Natural remanent magnetization intensities ranged from 7.6×10^{-3} to 1.4×10^{-1} A/m, although most of

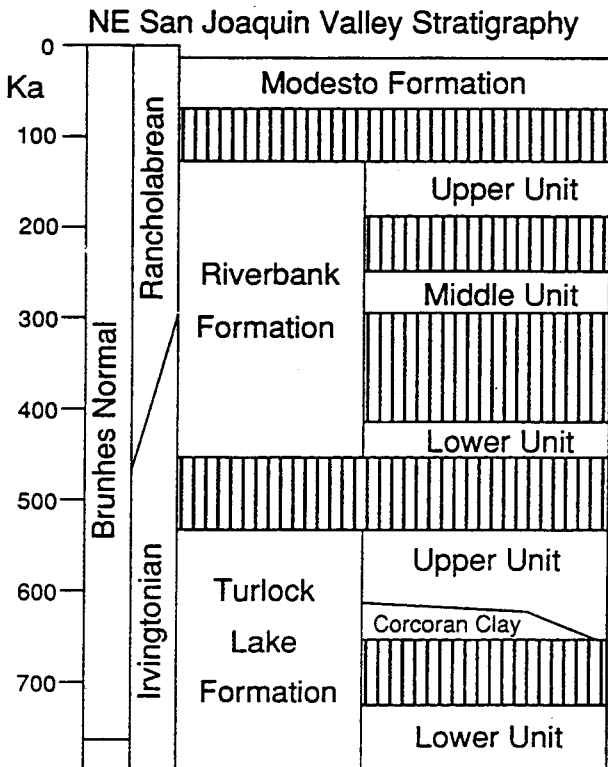


Figure 4. Younger Pleistocene stratigraphic units of the northeast San Joaquin Valley as defined by Marchand and Allwardt (1981).

the samples fell in the narrower range of 2.4×10^{-2} to 6.2×10^{-2} A/m. All of the samples were extremely well-behaved during the demagnetization procedure. The remanent magnetic directions of all of the samples decayed univectorially toward the origin. A few samples appeared to have minor secondary components that were removed after the first or second demagnetization step.

Declinations of the samples ranged from 30° W to 20° E. Inclinations ranged from 30° to 62° . These directions are fully consistent with a normal magnetic polarity. Because of the great stratigraphic sampling interval, it is unlikely that these normal polarities represent a short normal subchron in the Matuyama reversed chron. We conclude that the sediments at the Fairmead landfill, and in particular the fossil-bearing interval of Unit C, were deposited during the Brunhes normal polarity epoch. This places an upper bound of 780,000 years on the age of the fossils found at this site.

FAUNA

The vertebrate fossils recovered from Fairmead Landfill represent a diverse fauna dominated by large mammals. Microvertebrate material is rare at this site, but several tons of matrix remain to be processed which may yield additional small vertebrate taxa. The taxa identified thus far are listed in Table 1. These are tentative identifications; further study will permit assignment of some specimens to the species level. The preliminary identifications were made using comparative specimens in the research collections of the UCMP and pertinent literature on each of the taxa. Detailed descriptions of the faunal material are not provided here because much material is unprepared and the fauna is still under study. For reference, voucher specimens for each taxon are provided in Table 2. Selected specimens are illustrated in Plate I.

AGE

The age of the Fairmead Landfill fauna is late Irvingtonian, based on the currently known stratigraphic ranges of species present at the site (Figure 5) and from magnetostratigraphy and geologic data. Most of the taxa have stratigraphic ranges encompassing much of the Pleistocene. In part this is because many of the identifications are only to the genus level. Two species indicate an Irvingtonian age; *Tetrameryx irvingtonensis* and *Canis armbrusteri*. The presence of *Tetrameryx irvingtonensis* is suggestive of a middle-late Irvingtonian age because it is known only from the type Irvington locality in the San Francisco Bay area (Kurten and Anderson,

Table 1. Taxa identified from Fairmead Landfill, locality UCMP V93128.

Class Reptilia	Order Testudines	Family Emydidae	<i>Clemmys marmorata</i>
Class Aves	Order Anseriformes	Family Anatidae	
Class Mammalia	Order Edentata	Family Mylodontidae	<i>Glossotherium harlani</i>
		Family Megatheriidae	<i>Nothrotheriops</i> cf. <i>N. shastensis</i>
		Family Megalonychidae	<i>Megalonyx</i> sp.
	Order Carnivora	Family Canidae	<i>Canis</i> cf. <i>C. latrans</i>
			<i>Canis armbrusteri</i>
		Family Felidae	<i>Smilodon</i> cf. <i>S. fatalis</i>
			<i>Homotherium</i> sp.
	Order Rodentia	Family Geomyidae	<i>Thomomys</i> sp.
		Family Heteromyidae	cf. <i>Dipodomys</i> sp.
	Order Lagomorpha	Family Leporidae	<i>Lepus</i> sp.
	Order Proboscidea	Family Elephantidae	<i>Mammuthus columbi</i>
	Order Perissodactyla	Family Equidae	<i>Equus</i> sp.
	Order Artiodactyla	Family Camelidae	<i>Camelops</i> sp.
			<i>Hemiauchenia</i> sp.
		Family Antilocapridae	<i>Tetrameryx irvingtonensis</i>
			<i>Capromeryx</i> sp.
		Family Cervidae	<i>Odocoileus</i> sp.

1980). *Canis armbrusteri* is known from several Irvingtonian age localities in North America and the species became extinct around the Irvingtonian/Rancholabrean boundary (Kurten and Anderson, 1980; Nowak, 1979). Savage (1951) distinguished the Irvingtonian and Rancholabrean NALMAs on

Table 2. List of voucher specimens for taxa recovered from Fairmead Landfill, UCMP locality V93128. Abbreviations: R. = right, L. = left, i = lower incisor, dp = deciduous lower premolar, p = lower premolar, P = upper premolar, m = lower molar, M = upper molar.

<u>Taxon</u>	<u>Voucher specimen(s) with element identification</u>
<i>Clemmys marmorata</i>	UCMP 140625, incomplete carapace
Anatidae	UCMP 140414, distal humerus
<i>Glossotherium harlani</i>	UCMP 140263, premolar; UCMP 140392, R. dentary
<i>Nothrotheriops</i> cf. <i>N. shastensis</i>	UCMP 140260, cranium; UCMP 140261, tooth
<i>Megalonyx</i> sp.	UCMP 140262, upper canine; UCMP 140393 premolar or molar
<i>Canis arbrusteri</i>	UCMP 140265, partial R. dentary with p1-m2
<i>Canis</i> cf. <i>C. latrans</i>	UCMP 140413, L. maxilla fragment with M2-M3
<i>Smilodon</i> cf. <i>S. fatalis</i>	UCMP 140426, R. and L. frontal and parietal; UCMP 140264, distal fibula, partial L. astragalus
<i>Homotherium</i> sp.	UCMP 140390, partial cranium; UCMP 140391, upper canine
<i>Thomomys</i> sp.	UCMP 140408, lower premolar
cf. <i>Dipodomys</i> sp.	UCMP 140627, R. i1
<i>Lepus</i> sp.	UCMP 140626, incomplete juvenile R. tibia
<i>Mammuthus columbi</i>	UCMP 140415, R. and L. dentaries with teeth
<i>Equus</i> sp. (large and small)	UCMP 140404, L. dentary fragment with p2-p3
<i>Camelops</i> sp.	UCMP 140400, partial R. dentary with p4-m3; UCMP 140401, partial R. dentary with p4-m3
<i>Hemiauchenia</i> sp.	UCMP 140397, partial R. dentary with dp3-m2
<i>Tetrameryx irvingtonensis</i>	UCMP 140398, L. cranial fragment w. horn cores UCMP 140410, L. m2
<i>Capromeryx</i> sp.	UCMP 140409, radius
<i>Odocoileus</i> sp.	UCMP 140399, base of R. antler

the basis of the presence of *Bison* in Rancholabrean localities. The absence of *Bison* in a fauna represented by thousands of specimens, and dominated by large herbivores which normally occur with *Bison* in Rancholabrean localities, is further indicative of an Irvingtonian age.

Paleomagnetic analysis suggests a maximum age of 780,000 yr. B.P., the Matuyama/Brunhes boundary (Baksi et al., 1992), for the Fairmead Landfill fauna because the stratum in which the fossils occur is normally magnetized. As a great stratigraphic range was sampled, it is unlikely that the fossil-bearing unit was deposited during the Jaramillo normal subchron within the Matuyama reversed magnetic chron. The Jaramillo normal magnetic subchron and the beginning of the Brunhes normal magnetic chron are the only two major periods of normal magnetism that occur in the late Irvingtonian (Lindsay et al., 1987; Kurten and Anderson, 1980).

DISCUSSION

Savage (1951) originally defined the Irvingtonian NALMA based on the fauna from the gravel pits near Irvington, Alameda County, California. As presently construed, the Irvingtonian begins at about 1.9 Ma and ends between about 0.5-0.3 Ma (Lundelius et al., 1987). The Irvingtonian lacks long, continuous sequences of superimposed faunas such as those represented in the preceding Blancan NALMA. However, some short sequences exist (e.g. Vallecito Creek, San Diego County, California in part). Some localities have paleomagnetic data, while others have datable volcanic ash beds (Lundelius et al., 1987) but in large part Irvingtonian stratigraphy depends on fossil remains. Detailed study of the Fairmead Landfill fauna and revision of the type Irvington fauna should yield a picture of the middle to late Irvingtonian of central California. This will allow for better comparison and correlation with other Irvingtonian faunas in North America, which is the essence of biostratigraphy.

Taxa	Irvingtonian	Rancholabrean	Holocene
<i>Clemmys marmorata</i>			
Anatidae			
<i>Glossotherium harlani</i>			
<i>Nothrotheriops</i> cf. <i>N. shastensis</i>			
<i>Megalonyx</i> sp.			
<i>Canis armbrusteri</i>			
<i>Canis</i> cf. <i>C. latrans</i>	—		
<i>Smilodon</i> cf. <i>S. fatalis</i>	—		
<i>Homotherium</i> sp.			
<i>Thomomys</i> sp.		—	
cf. <i>Dipodomys</i> sp.			
<i>Lepus</i> sp.			
<i>Mammuthus columbi</i>	—		
<i>Equus</i> sp.			
<i>Camelops</i> sp.			
<i>Hemiauchenia</i> sp.			
<i>Tetrameryx irvingtonensis</i>	—		
<i>Capromeryx</i> sp.			
<i>Odocoileus</i> sp.			

Figure 5. Stratigraphic ranges of Fairmead Landfill taxa. Ranges were compiled from Kurtén and Anderson (1980) and Nowak (1979).

ACKNOWLEDGEMENTS

We express our appreciation to the County of Madera, to Madera Disposal Systems, Inc. and especially to the many volunteers for their time and assistance with excavation. Diane L. Blades deserves special recognition for her efforts, initially as a volunteer, and later as on-site monitor for the excavation from September 1993 to the present time. Volunteers of special note who have contributed significantly to the excavation include: Bill

Hamilton, Sam Champion, Wally Nichols, Pam Lara, Mary DiViccaro and John McCormick. Thanks to Art Alejandre, formerly of Madera Disposal Systems, for his assistance throughout 1993 and 1994. The faunal remains were identified by R. G. Dundas, J. H. Hutchison, C. J. Bell and L. G. Nelms. The 1993 UCMP fossil recovery work was supervised by L. G. Nelms, J. H. Hutchison and R. G. Dundas. Subsequent fossil recovery work has been conducted under the direction of R. G.

Plate I. Each specimen illustrated in Plate I has its own bar scale. The bar scale following each letter on the plate is equal to 2 cm.

Figure A. *Canis armbrusteri*, UCMP 140265, a partial right dentary with p1-m2. Lateral view.

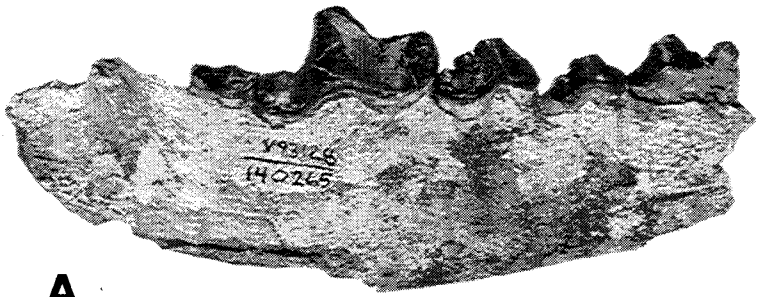
Figure B. *Odocoileus* sp., UCMP 140399, base of a right antler. Medial view.

Figure C. *Tetrameryx irvingtonensis*, UCMP 140398, left cranial fragment with partial horn cores.

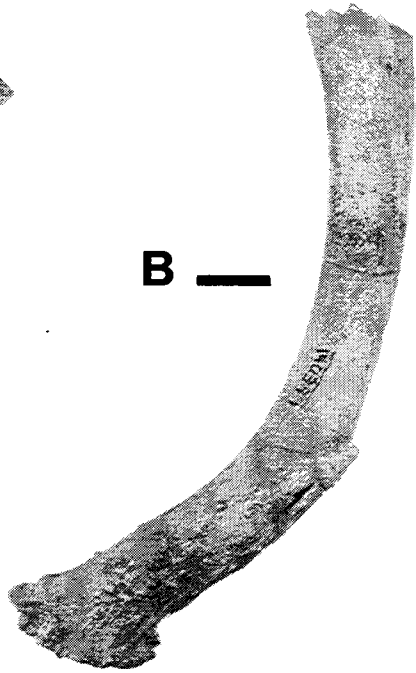
Figure D. *Camelops* sp., UCMP 140401, a partial right dentary with p4-m3. Lateral view.

Figure E. *Homotherium* sp., UCMP 140390, a partial cranium. Lateral view.

Figure F. *Homotherium* sp., UCMP 140391, upper canine. Lateral view.



A —



B —



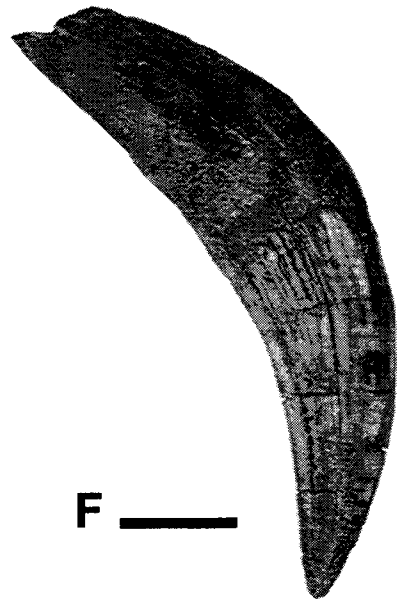
C —



D —



E —



F —

Dundas, J. H. Hutchison, D. E. Savage and C. J. Bell provided helpful reviews of the manuscript.

LITERATURE CITED

- Baksi, A. K., V. Hsu, M. O. McWilliams, and E. Farrar. 1992. $^{40}\text{Ar}/^{39}\text{Ar}$ Dating of the Brunhes-Matuyama Geomagnetic Field Reversal. *Science* 256:356-357.
- Fay, L. P. and K. R. Thiessen. 1993. A Preliminary Note on the Arvin Landfill Local Fauna (?Holocene), Kern County, California. *In* New Additions to the Pleistocene Vertebrate Record of California (R. G. Dundas and D. J. Long, eds). *PaleoBios* 15(4):78-80.
- Firby, J. B. 1968. Revision of the Middle Pleistocene Irvington Fauna of California. Unpublished M.A. Thesis, University of California, Berkeley. 134 pp.
- Helley, E. J. 1966. Sediment Transport in the Chowchilla River Basin: Mariposa, Madera, and Merced Counties, California. Unpublished Ph.D. Dissertation, University of California, Berkeley. 153 pp.
- Kurtén, B. and E. Anderson. 1980. Pleistocene Mammals of North America. Columbia University Press, New York. 442 pp.
- Lindsay, E. H., N. M. Johnson, and N. D. Opdyke. 1975. Preliminary correlation of North American Land Mammal Ages and geomagnetic chronology. Pages 111-119 *in* Studies on Cenozoic Paleontology and Stratigraphy (G. R. Smith and N. E. Friedland, eds.). University of Michigan Papers on Paleontology 12.
- Lindsay, E. H., N. M. Johnson, N. D. Opdyke, and R. F. Butler. 1987. Mammalian Chronology and the Magnetic Polarity Time Scale. Pages 269-284 *in* Cenozoic Mammals of North America: Geochronology and Biostratigraphy (M. O. Woodburne, ed.). University of California Press, Berkeley.
- Lundelius, E. L., Jr., T. Downs, E. H. Lindsay, H. A. Semken, R. J. Zakrzewski, C. S. Churcher, C. R. Harington, G. E. Schultz, and S. D. Webb. 1987. The North American Quaternary sequence. Pages 211-235 *in* Cenozoic Mammals of North America: Geochronology and Biostratigraphy (M. O. Woodburne, ed.). University of California Press, Berkeley.
- Marchand, D. E. and A. Allwardt. 1981. Late Cenozoic Stratigraphic Units, Northeastern San Joaquin Valley, California. *United States Geological Survey Bulletin* 1470:1-70.
- Nowak, R. M. 1979. North American Quaternary *Canis*. Monograph of the Museum of Natural History, University of Kansas 6:1-154.
- Savage, D. E. 1951. Late Cenozoic Vertebrates of the San Francisco Bay Region. University of California Publications, *Bulletin of the Department of Geological Sciences* 28:215-314.

THE FAIRMEAD LANDFILL LOCALITY (PLEISTOCENE, IRVINGTONIAN), MADERA COUNTY, CALIFORNIA

DUNDAS, Robert G., Department of Geology, California State University, Fresno, CA 93740, rdundas@csufresno.edu; BLADES, Diane L., San Joaquin Valley Paleontology Foundation, P.O. Box 849, Chowchilla, CA 93610

A diverse fauna, dominated by terrestrial mammals, was discovered in May of 1993 at the Madera County Fairmead Landfill in alluvial fan, fan channel and marsh/lacustrine sediments representing the upper unit of the Turlock Lake Formation. Taxa identified from the fauna include: Caudata, Anura, Colubridae, *Clemmy marmorata*, Anatidae, *Paramylodon harlani*, *Nothrotheriops* cf. *N. shastensis*, *Megalonyx* sp., *Canis armbrusteri*, *Canis* cf. *C. latrans*, *Vulpes* sp., *Smilodon* cf. *S. fatalis*, *Homotherium* sp., *Arctodus* sp., *Taxidea taxus*, *Thomomys* sp., *Spermophilus* sp., *Neotoma* sp., cf. *Dipodomys* sp., *Peromyscus* sp., *Microtus* sp., *Lepus* sp., *Mammuthus* sp., *Equus* sp. (large and small species), *Platygonus* sp., *Camelops* sp., *Hemiauchenia* sp., *Tetrameryx irvingtonensis*, *Capromeryx* sp., and *Odocoileus* sp. A late Irvingtonian age is indicated for the fauna based largely on the presence of *Tetrameryx irvingtonensis* coupled with the absence of *Bison*. Excavation at the site continues as part of the mitigation procedures recommended to the County of Madera in order to comply with California Environmental Quality Act (CEQA) regulations. The Fairmead Landfill is the first diverse late Irvingtonian fauna discovered in north-central California. Comparison to the older, coastal type Irvingtonian fauna will enhance our understanding of the Irvingtonian of California, permitting better comparisons and correlations with other North American Irvingtonian age faunas.

Geological Society of America, Abstracts with Programs, v. 31, n. 7.

The Fairmead Landfill Fossil Site: Continuing Recovery of a Diverse Middle Pleistocene (Late Irvingtonian) Biota in Madera County, California

Niranjala Kottachchi, Robert G. Dundas and James C. Chatters, Department of Earth & Environmental Sciences, California State University, Fresno, CA 93740

Situated on the alluvial fan of the Chowchilla River, the Fairmead Landfill locality represents the largest middle Pleistocene biota in the San Joaquin Valley of California. Discovered in May of 1993, during excavation of a new expansion cell, the site has produced a diverse assemblage dominated by equids, camelids, and proboscideans. Fossils occur as isolated elements and in bone beds, preserved in deposits representing alluvial fan, fan channel, and marsh/lacustrine sediments of the upper unit of the Turlock Lake Formation. Over 5000 fossil specimens belonging to at least three dozen taxa have been discovered at the site, including: Caudata, Anura, Colubridae, *Clemmys marmorata*, *Xerobates agassizi*, Anatidae, *Paramylodon harlani*, *Nothrotheriops cf. texanus*, *Megalonyx wheatleyi*, *Notiosorex*, *Canis* (wolf), *Canis cf. C. latrans*, *Vulpes*, *Smilodon*, *Homotherium*, *Arctodus*, *Taxidea taxus*, *Thomomys*, *Spermophilus*, *Neotoma*, cf. *Dipodomys*, *Peromyscus*, *Microtus*, *Lepus*, *Mammuthus columbi*, *Equus*, *Platygonus*, *Camelops*, *Hemiauchenia*, *Tetrameryx irvingtonensis*, *Capromeryx*, and *Odocoileus*, as well as fossil plants, gastropods and bivalves. Based primarily on the presence of *Tetrameryx irvingtonensis* and the absence of *Bison*, a late Irvingtonian age is assigned to the locality.

To date, fossils have been collected from an area of over 14 acres. It is anticipated that paleontological monitoring and fossil salvage will continue for the duration of landfill activities, another 20+ years.

CSU Fresno assumed responsibility for the paleontological recovery and preparation in September 2007 and now collaborates with the San Joaquin Valley Paleontological Foundation to exploit this unique window into the Middle Pleistocene of Central California. The extensive collection is being prepared and identified to make it available for scientific study and public education in a museum planned by Madera County.

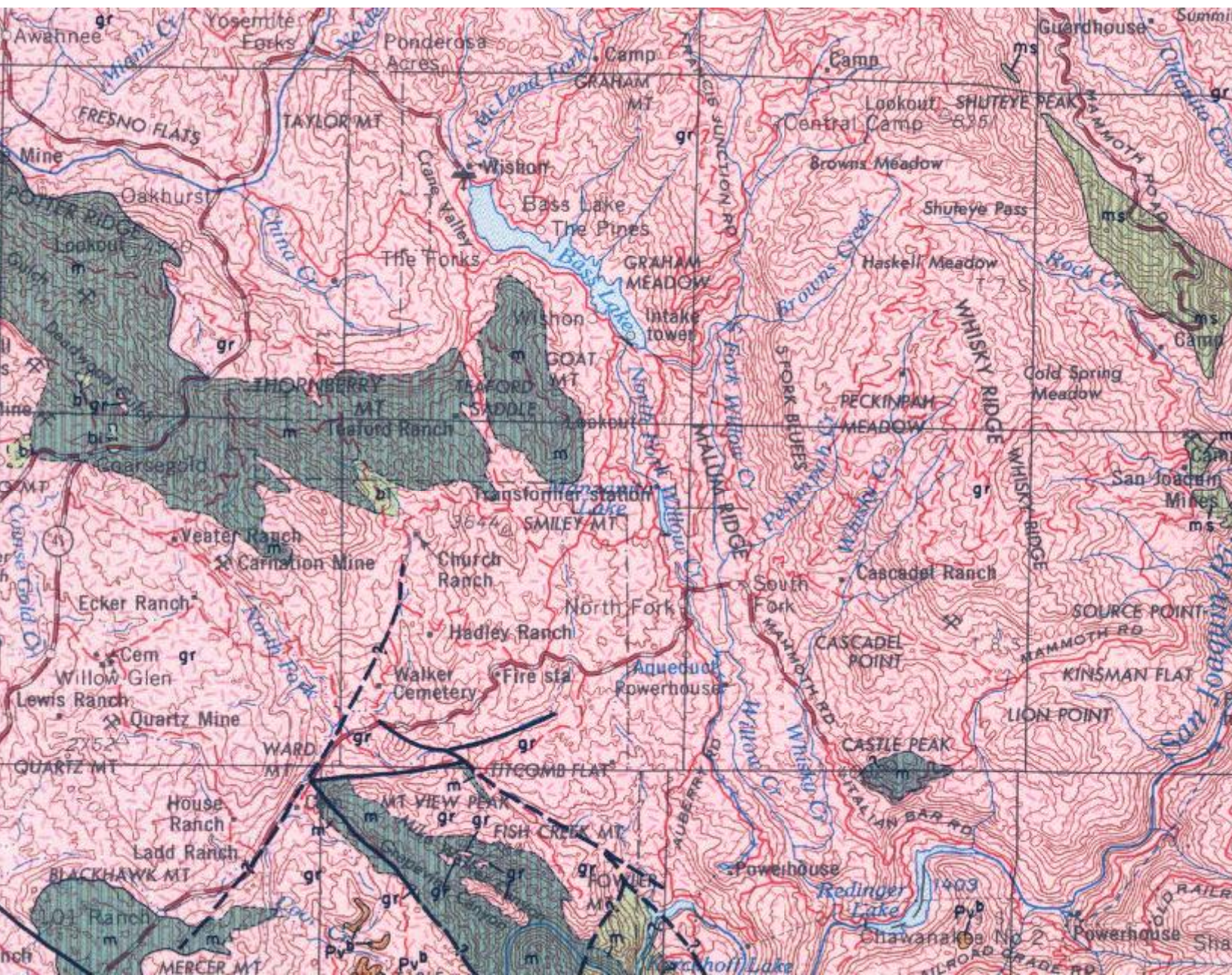


Plate I: North Fork Site locality depicting Mesozoic granites (section from Strand, 1967: Geologic map of California: Mariposa Sheet, scale 1:250,000)



Plate II: Views of the North Fork Site dominated by Mesozoic granitic igneous rocks

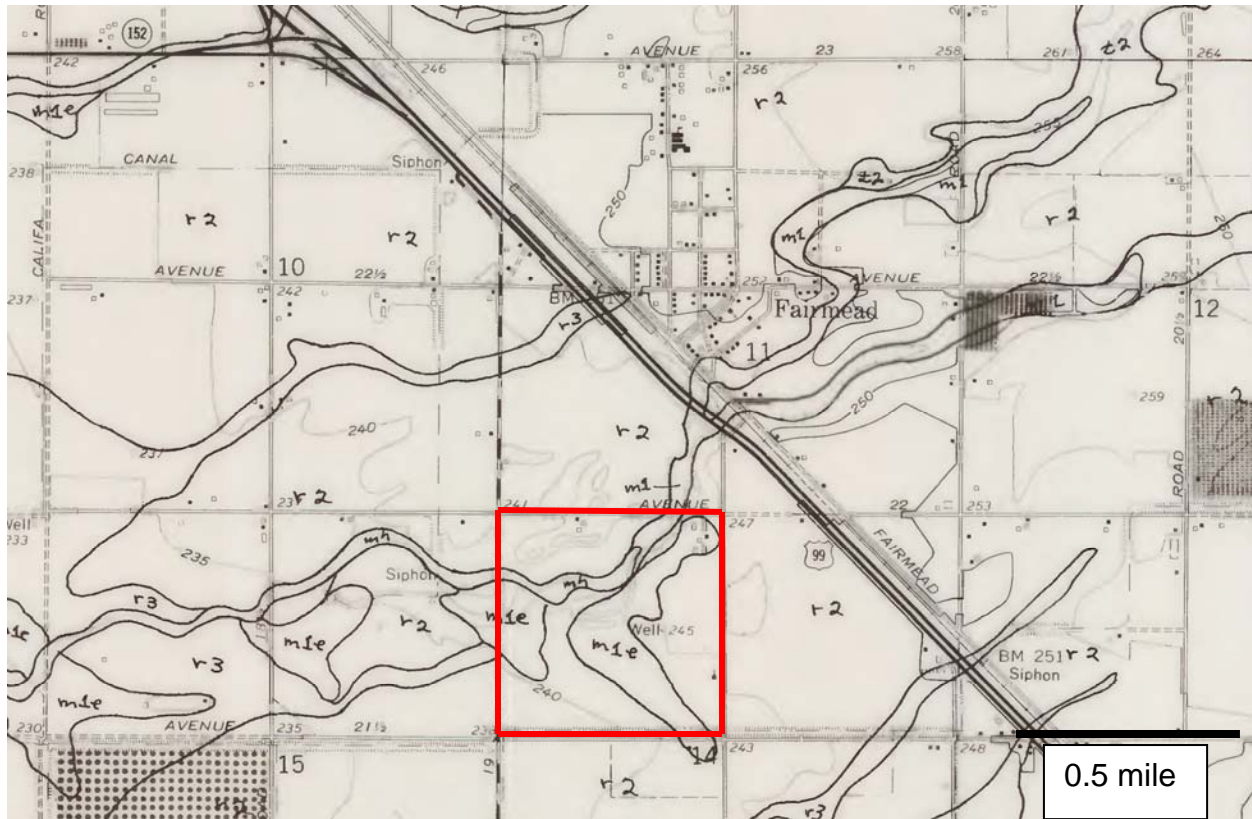


Plate III: Preliminary Quaternary Geologic Map of the Chowchilla Area, California (Marchand, 1976; section extracted from sheet 4 of 5, open file report 76-839). Location of Fairmead Landfill is enclosed within red box.

mh = undifferentiated Modesto and Holocene

m1e = lower member of Modesto Formation

r2 = middle member of Riverbank Formation

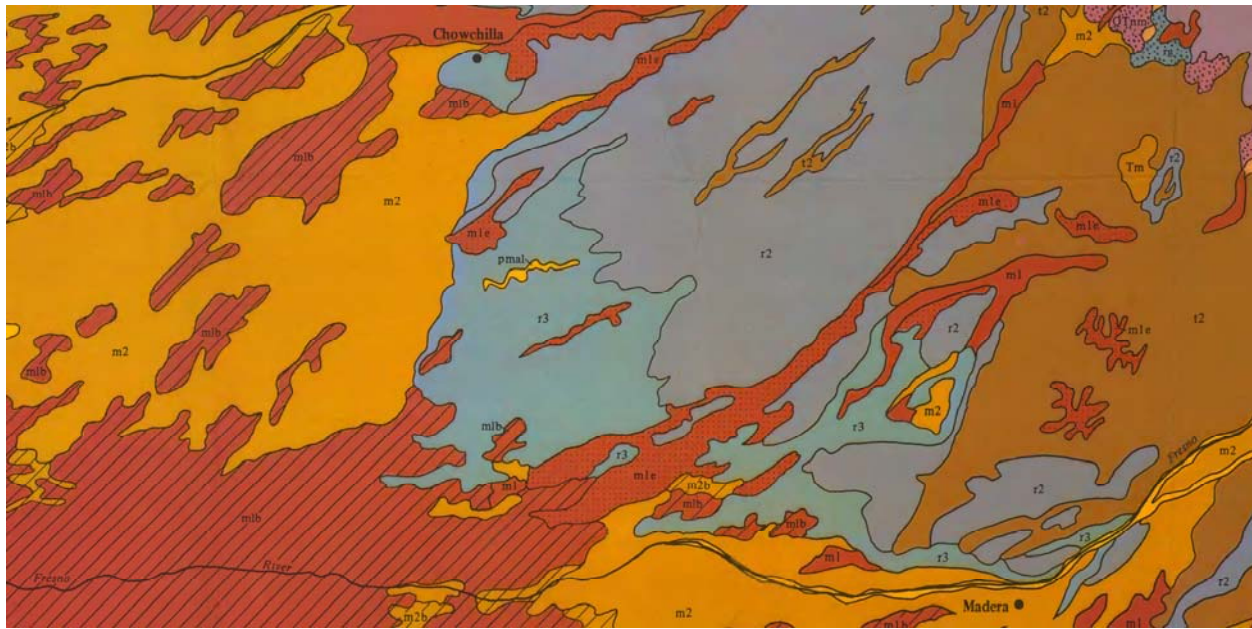
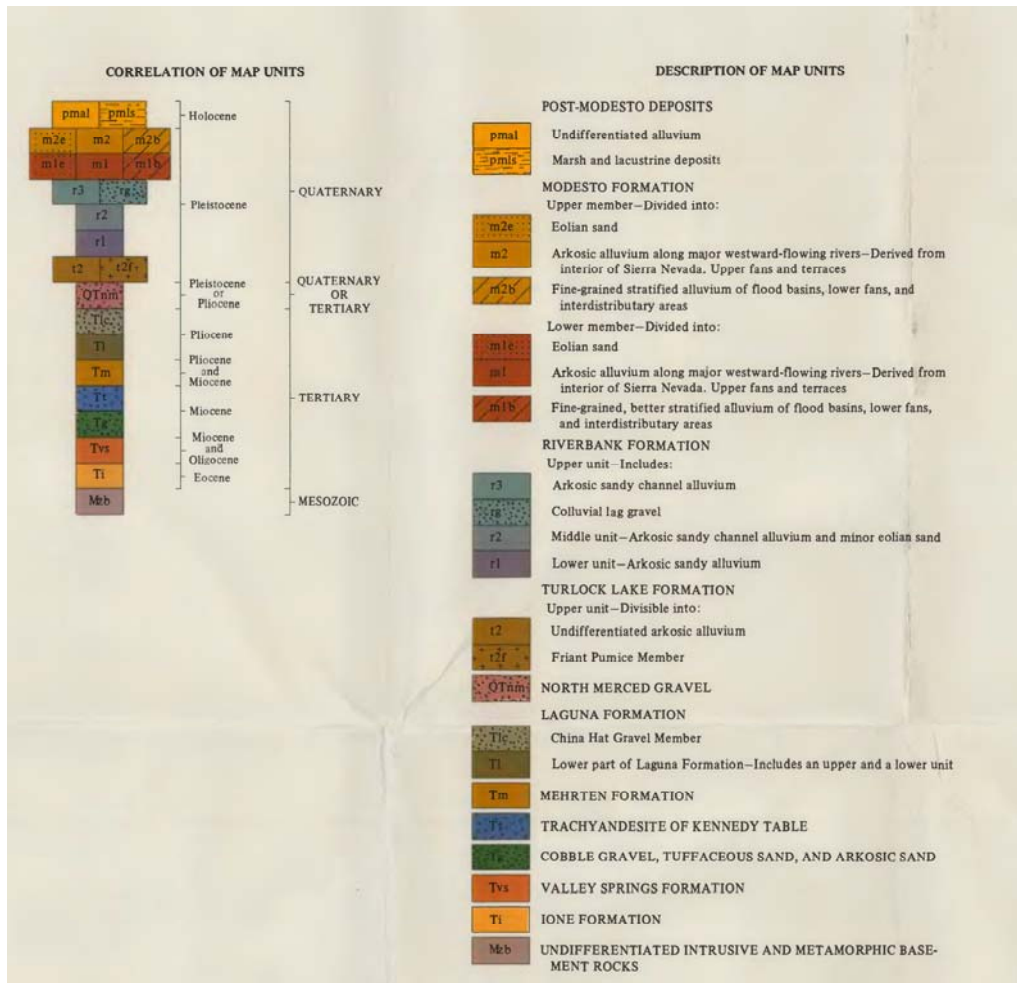


Plate IV: Portion of the generalized geologic map of the Merced-Madera Area depicting stratigraphic units present at both the Fairmead Landfill Site and study area (Marchand and Allwardt, 1981) 2 miles



Legend to accompany Plate IV



Plate V: Views of the Madera Site location